

# Indian Ocean and South-Asian monsoon: upper-ocean processes relevant to monsoon annual cycle

**H. Annamalai<sup>1</sup>, Motoki Nagura<sup>2</sup>, Kelvin Richards<sup>1</sup>, Caroline Ummenhofer<sup>3</sup>**

1 International Pacific Research Center (IPRC)/Department of Oceanography, Univ. Hawaii, USA

2 JAMSTEC, Japan

3 Woods Hole Oceanographic Institution, Woods Hole, USA

**“A big thanks to Jay”**

**“Thanks to reviewers on the first draft”**

**IndOOS Review Meeting (March 22-23), Jakarta**

# Talk Outline

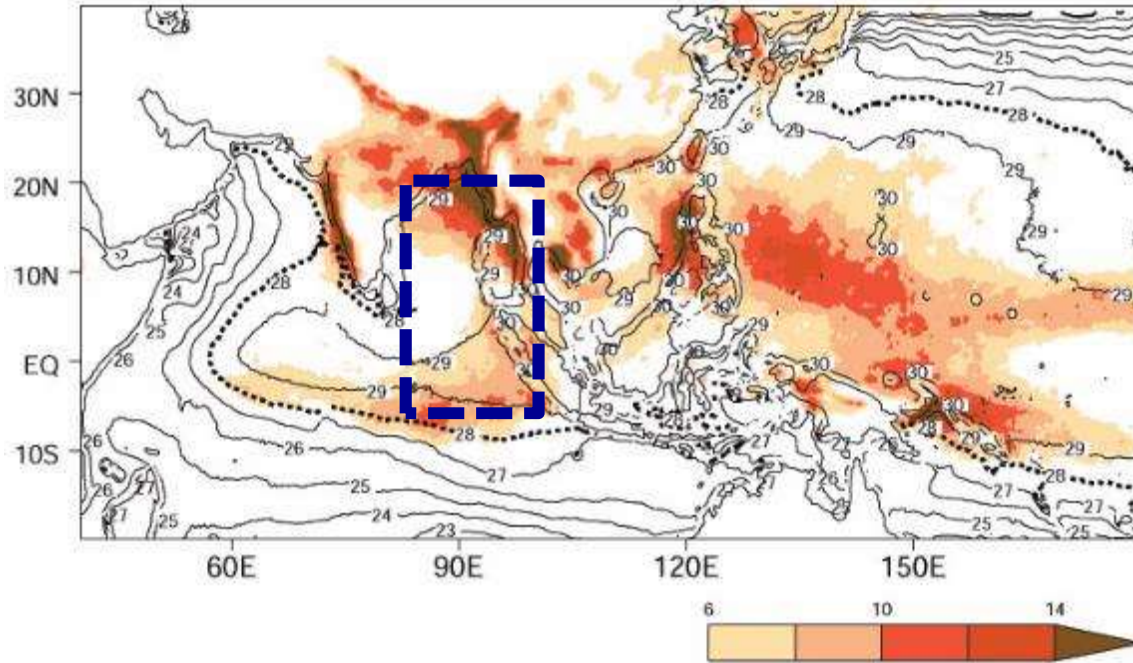
- Model systematic errors in ocean-atmosphere variables persist for the last few decades (despite data available from multiple observational platforms (in-situ, satellite, field experiments))
- Model errors - limitations in representing multiple processes and their interactions
- Current suite of observations not adequate enough to **understand the monsoons and constrain model physics**

Focus on upper-ocean processes (**thermal and salinity stratification**)

(Shenoi et al. 2002; Schott et al. 2009; Seo et al. 2009; Annamalai et al. 2015; Thadathil et al. 2016; Nagura et al. 2015; Mahadevan et al. 2016)

# Precipitation and SST

## Boreal Summer (JJA)



EIO SST remains  $> 28^{\circ}\text{C}$

Equatorial, eastward Wyrтки Jets

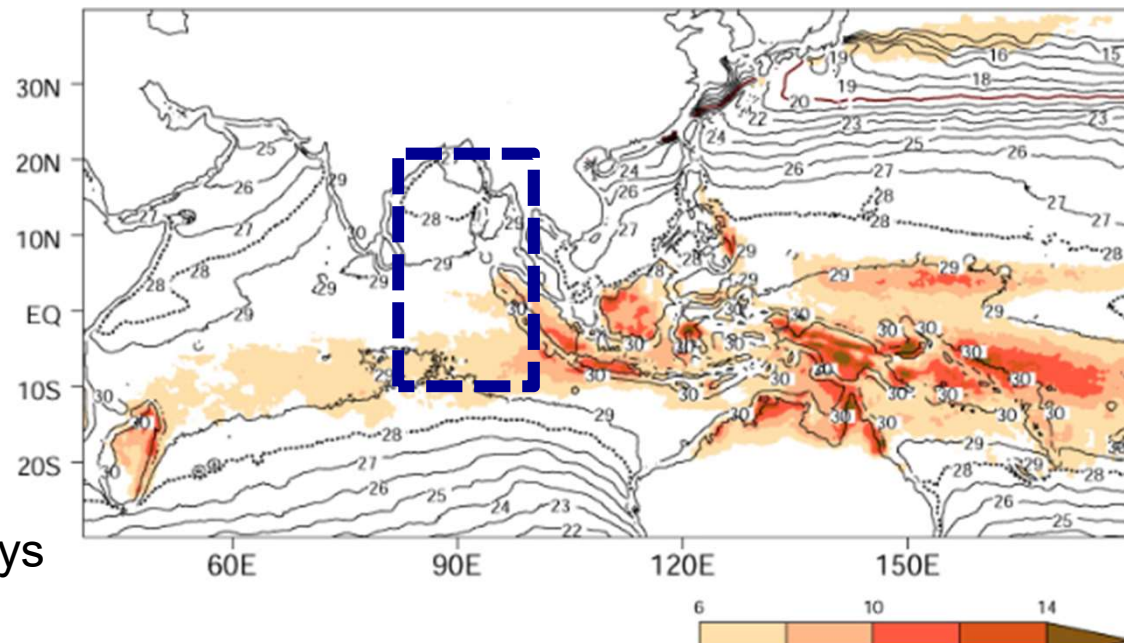
Precip occurs year-around  
Upper-ocean stratification

Bay of Bengal SST  $> 28^{\circ}\text{C}$

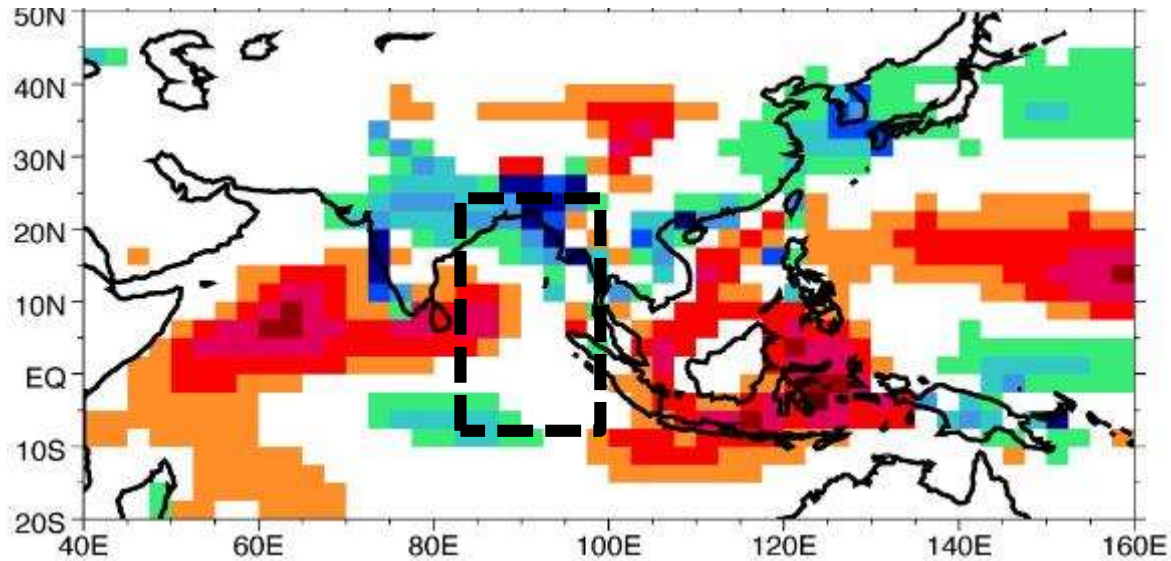
Upper-ocean stratification

Cold/dry winter monsoon  
near-surface air temperature  
drops  $< 26^{\circ}\text{C}$  but SST remains  
High – Temperature inversion  
peaks (Nov-Dec) – RAMA buoys  
(Thadathi I et al. 2006)

## Boreal Winter (DJF)



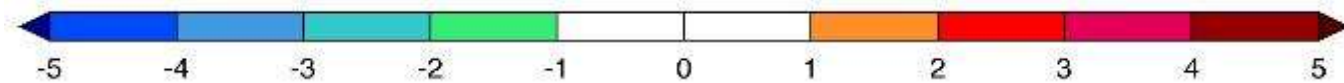
## CMIP3 MMM – GPCP



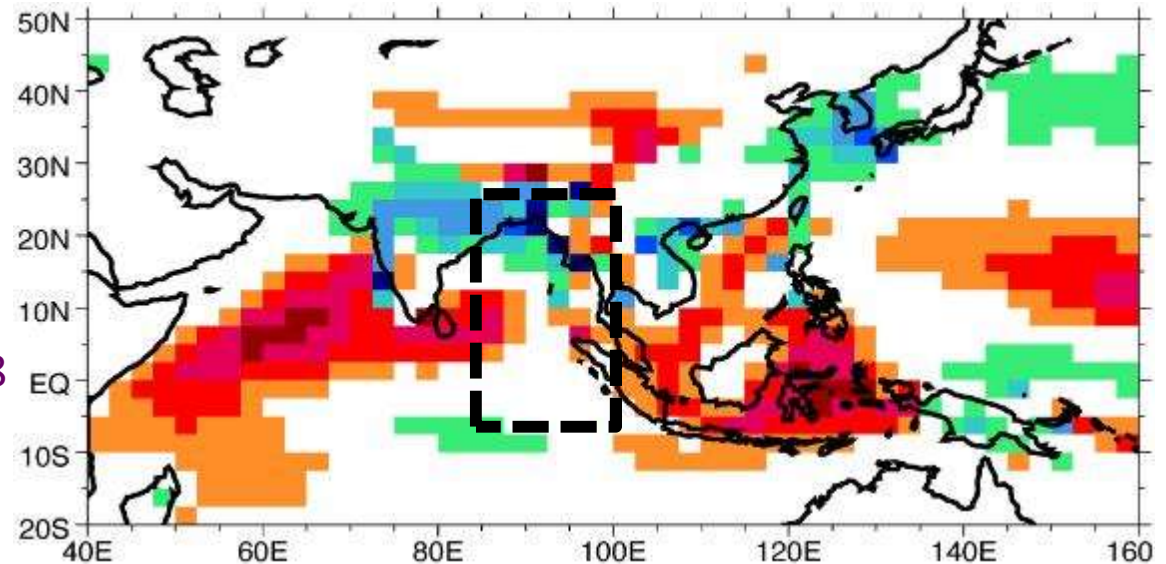
**JJAS - Precipitation**

(Sperber et al. 2013)

(mm/day)



## CMIP5 MMM – GPCP

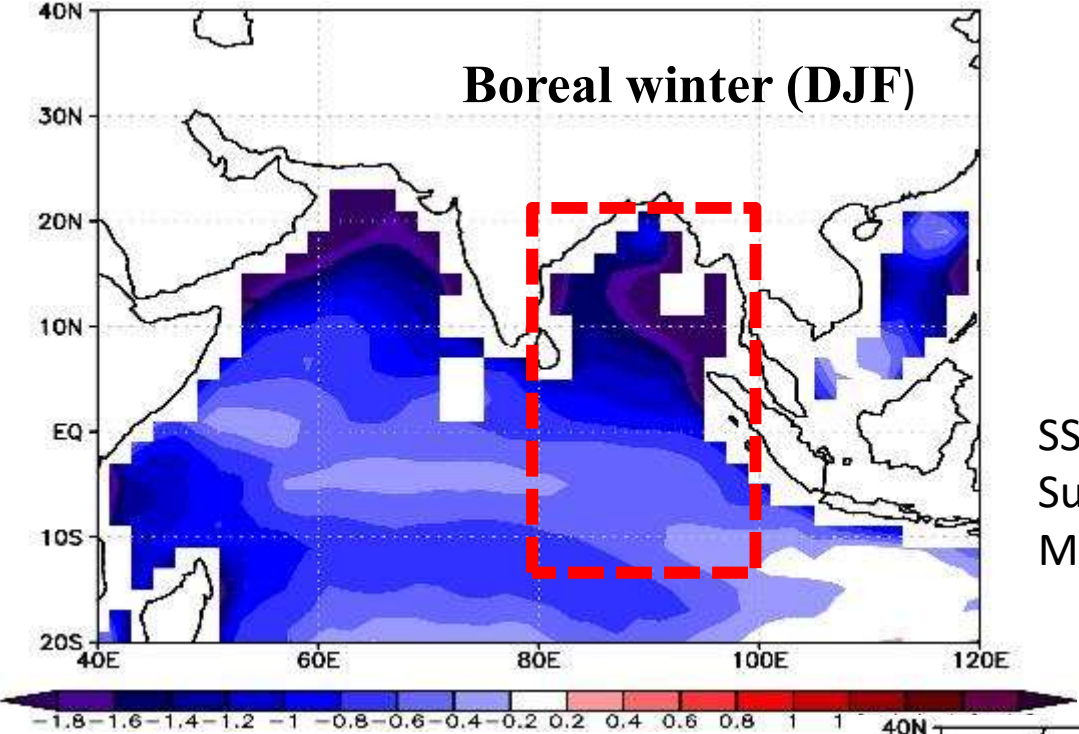


**Bay of Bengal  
Eastern EIO**

Nagura, McCreary and Annamalai 2018

BoB precip and salinity errors-  
EICC – stratification in Northern AS  
Errors in Mixed-layer depth and D20

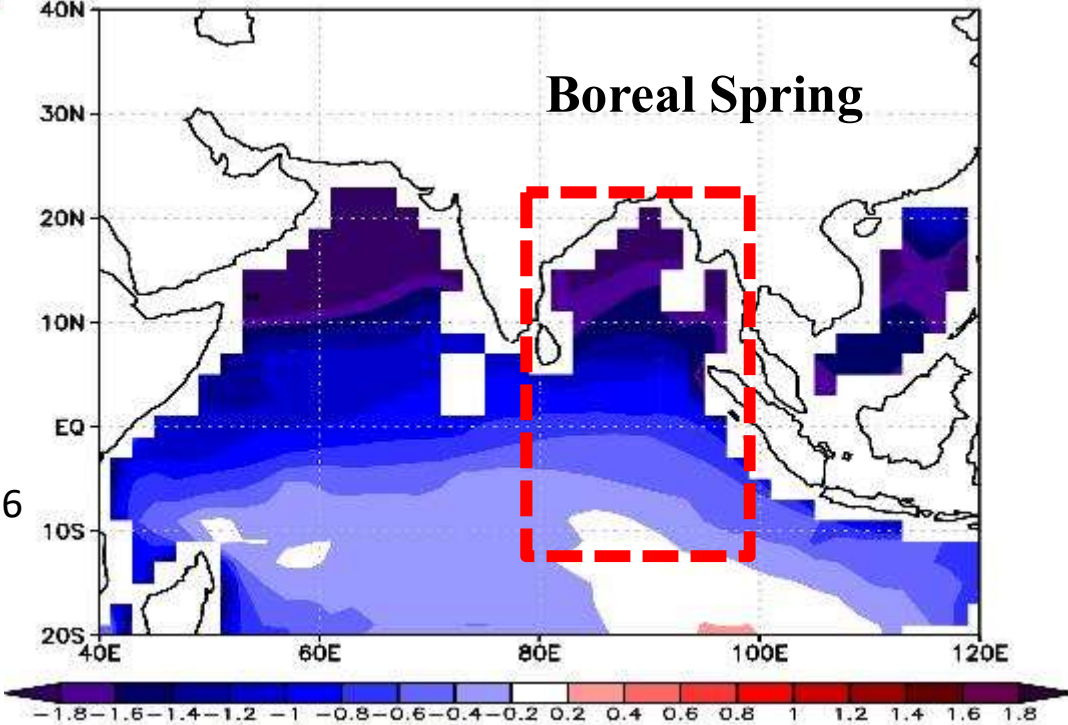
# CMIP5 MMM SST biases



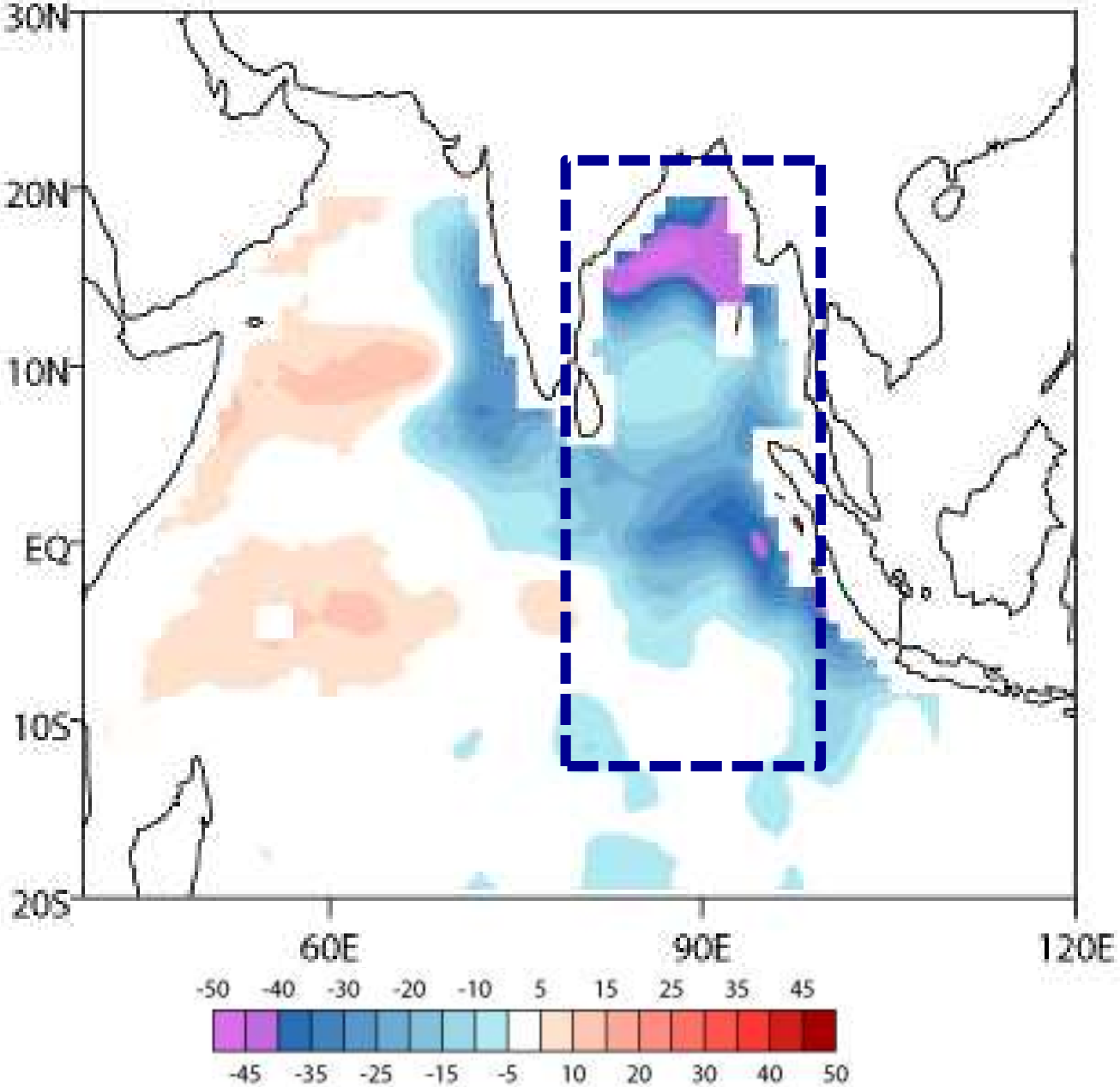
SST biases persist into the boreal spring and Summer seasons impacting the ensuing Monsoon during summer

“ SST – precip – SST “ In regions of high mean precip, two-way interaction needs to be understood and modeled

Levin et al. 2013; Sandeep and Ajayamohan 2016  
Annamalai 2018



# CMIP5 MMM errors in barrier-layer thickness (DJF)

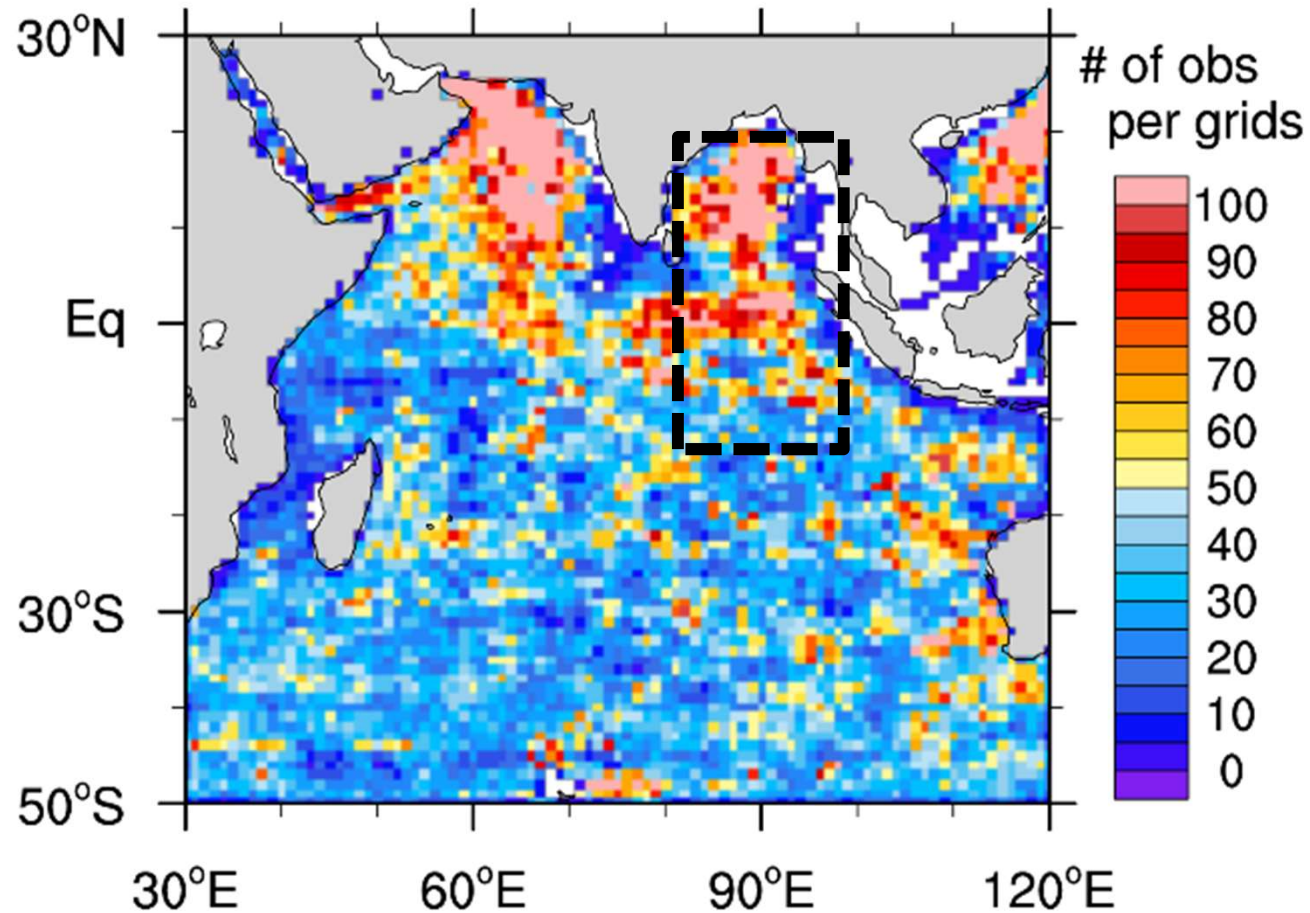


“SST-> Pr -> SST”

Salinity and temperature  
Stratifications in the  
Upper-ocean

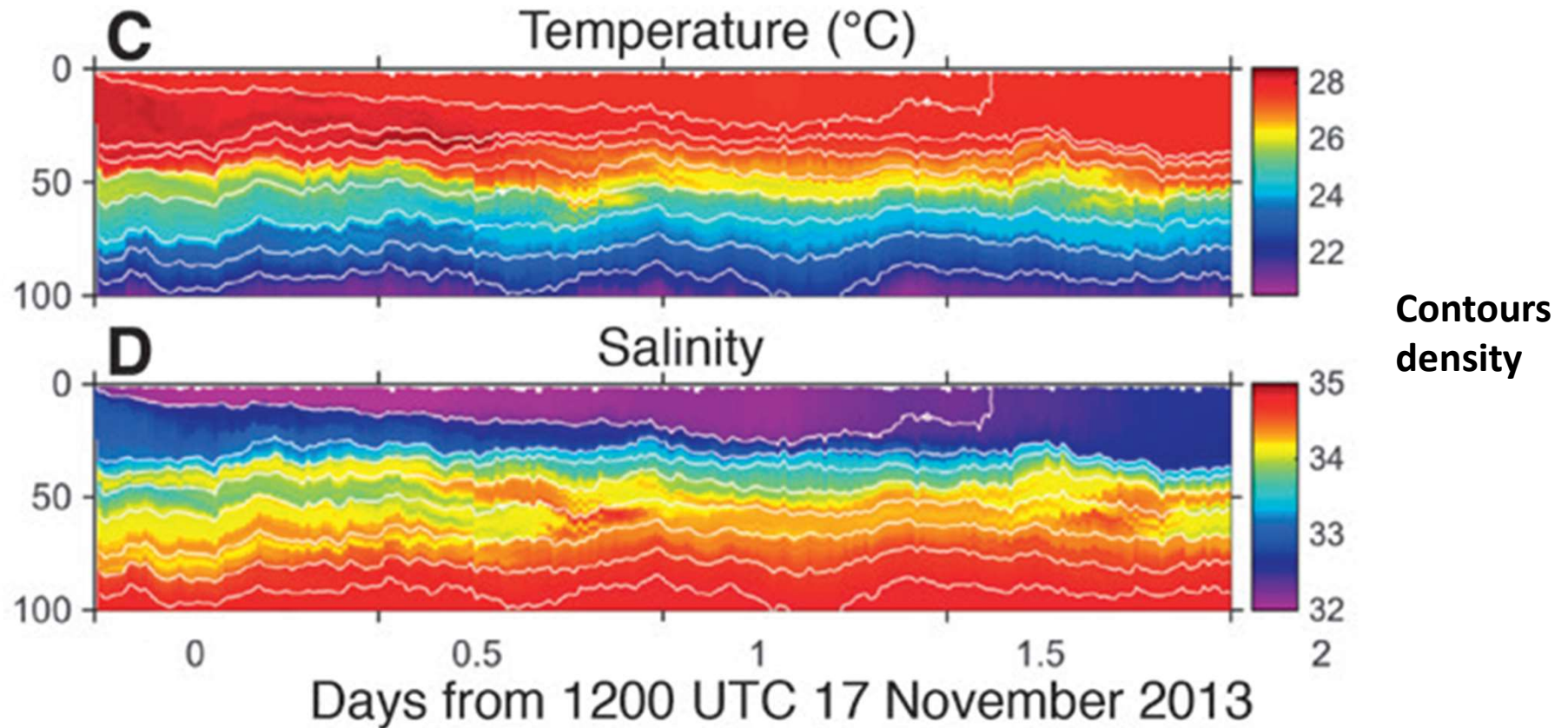
Mixed-layer < 10 m

## Number of in-situ observations (CTD/Argo floats) – Since 1961 To measure temperature and salinity



1. Argo floats – limitations (top 25 m – where stratification is strong)
2. Eastern EIO / coastal regions (limited observations)
3. Limited observations over the SIO Thermocline ridge / Arabian Sea warm pool

## An Ocean-Atmosphere Initiative for Bay of Bengal (ASIRI)



**A 2-day drift of two wire walkers near 16°N, 87°E**

Wijesekera et al. (2016)

Anderson and Riser (2014)  
Argo – auxiliary STS sensors  
Vertical res. 10 cm (top 2-4 m)  
Stratification / ver. gradient

Ocean-model vertical resolution needed? Kelvin Richards



## Recent/on-going observations field camps relevant to Upper-ocean processes

- **An Ocean-Atmosphere Initiative for Bay of Bengal (ASIRI) –**

High-resolution observations from multiple platforms

Shallow salinity-controlled mixed-layers result from high rainfall and river-runoff

Mixed-layers over the northern Bay are of the  $\sim 20\text{m}$ , and meso-scale oceanic features contribute to upper-ocean stratification (Nov-Dec 2013)

- **Auxiliary surface temperature and salinity (STS) sensors – Argo floats**

$\sim 10\text{ cm}$  vertical resolution in the top 15 m / strong vertical gradient (3 sites – BoB)

- **SPURS-2 Salinity Processes in the Upper-ocean Regional Study**

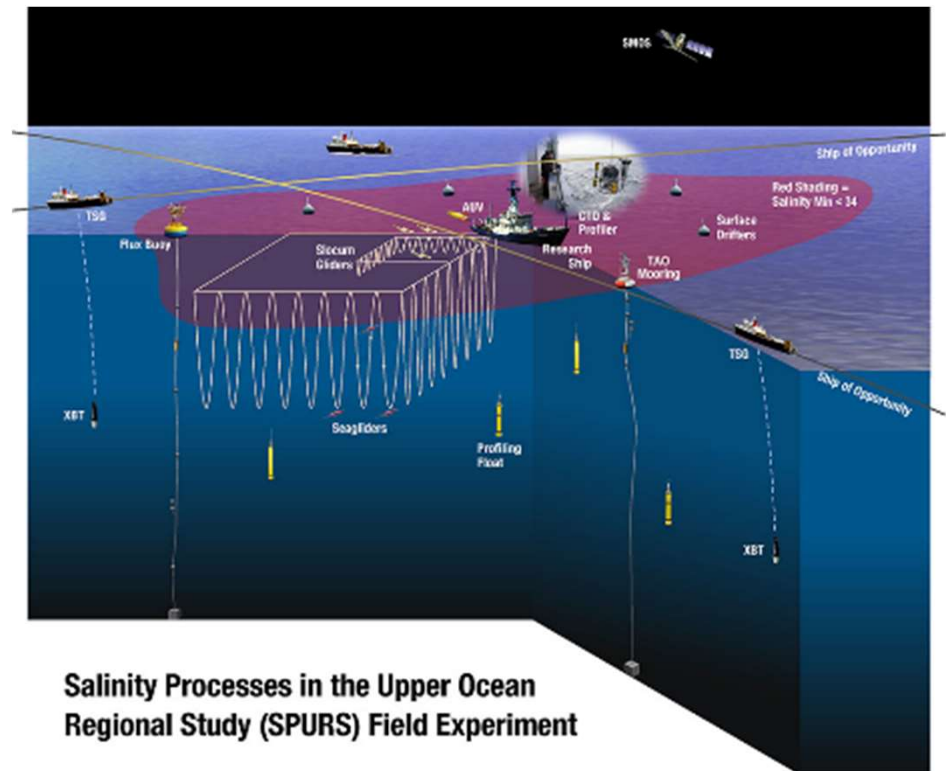
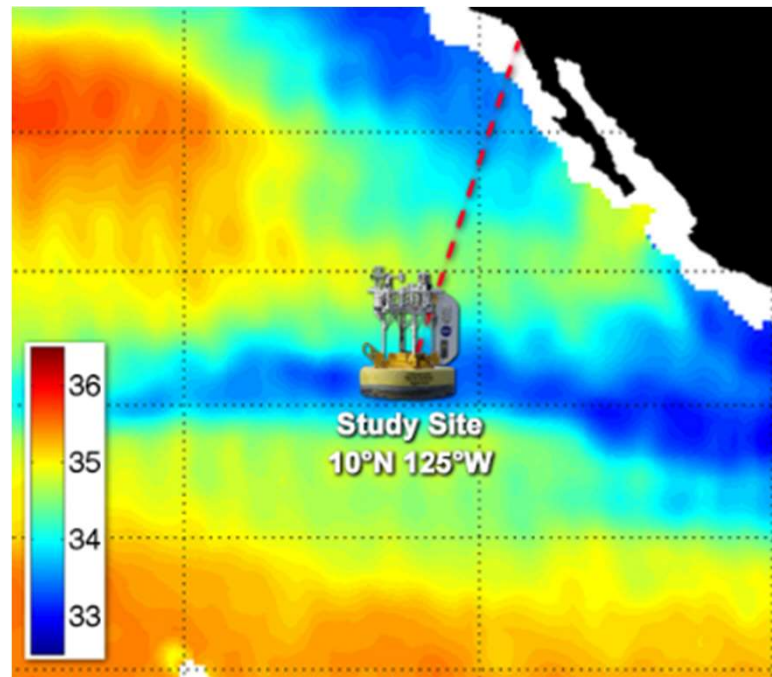
Eastern Pacific ITCZ (high-mean precipitation region)

Wijesekera et al. (2016); Anderson and Riser (2014); Rainville et al. (2017); Lindstorm et al. (2017 –

# Multiple platforms observations for sustained monitoring

## Upper-ocean stratifications

In regions of high-mean precipitation such as BoB/EEIO



# Actionable Recommendations of EOV

## Upper-ocean processes – Bay of Bengal – 12°-22°N, 80°-100°E

Temperature and Salinity – upper ocean; near-surface atmospheric variables

1. Moored buoys (2°x2° degree or even finer)
2. Argo floats – equip with auxiliary STS sensors (~10 cm ver. res top 2-20m)
3. Repeated glider sections and a suite of autonomous vehicles as in SPURS – vehicles such as SVP and SVP-2 drifters, sea/wave gliders,
4. SAILDRONES (being tested now along the equatorial western Pacific)

Rectangular mesh: Estimation of advection of salinity in and out of Bay of Bengal  
To augment reliability and in regions of shallow mixed-layers, repeated glider sections (Rainville et al. 2017) and a suite of autonomous vehicles (Lindstorm et al. 2017)

# SAILDRONES ARE USVs DESIGNED FOR **LONG RANGE, LONG DURATION** AUTONOMOUS MISSIONS

wind power for propulsion

solar power for electronics

20 feet tall

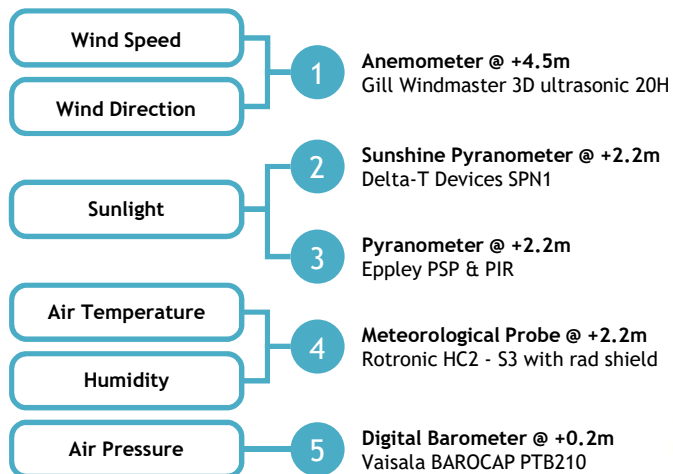
satellite link for live data

23 feet long

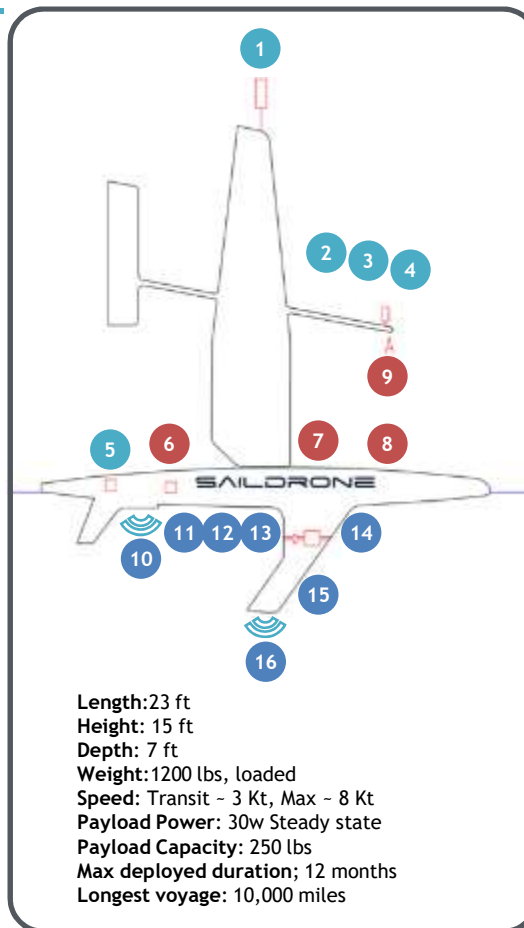
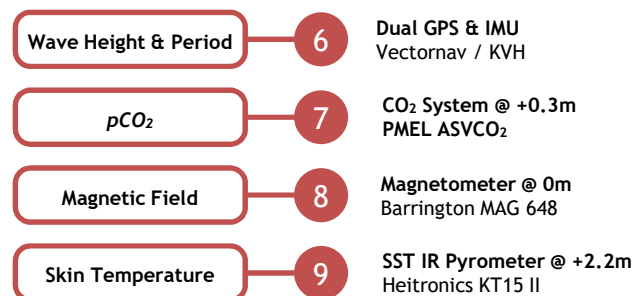


# COLLECTING AND TRANSMITTING REAL-TIME DATA

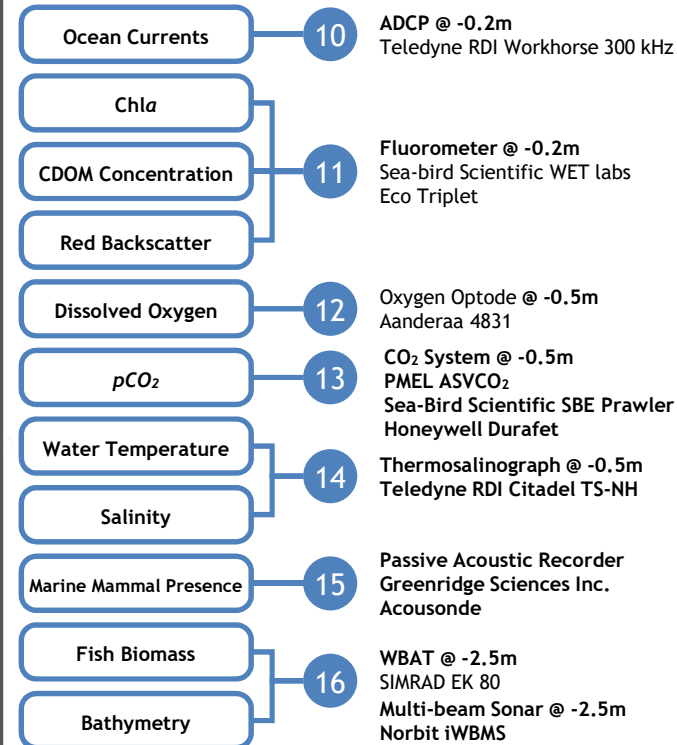
## Atmospheric Measurements



## Oceanic Surface Measurements



## Oceanic Sub-Surface Measurements





## SAILDRONES ARE SCALABLE IN BOTH FLEET SIZE AND VEHICLE SIZE

### SAILDRONE MAXI

- 72' long
- 55' tall wing
- 1 Ton payload
- 10-15 knots cruise

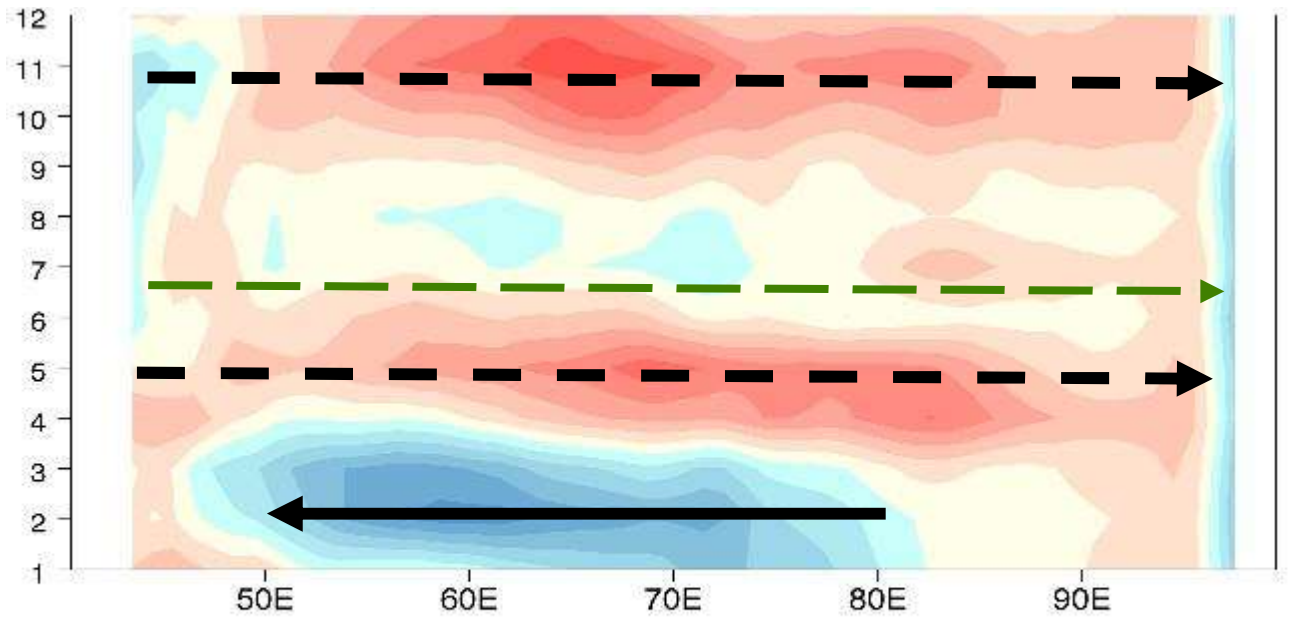




contact: [info@saildrone.com](mailto:info@saildrone.com)



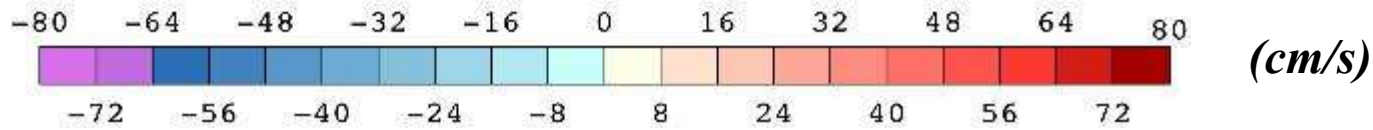
SAILDRONE



3°S-3°N – surface currents

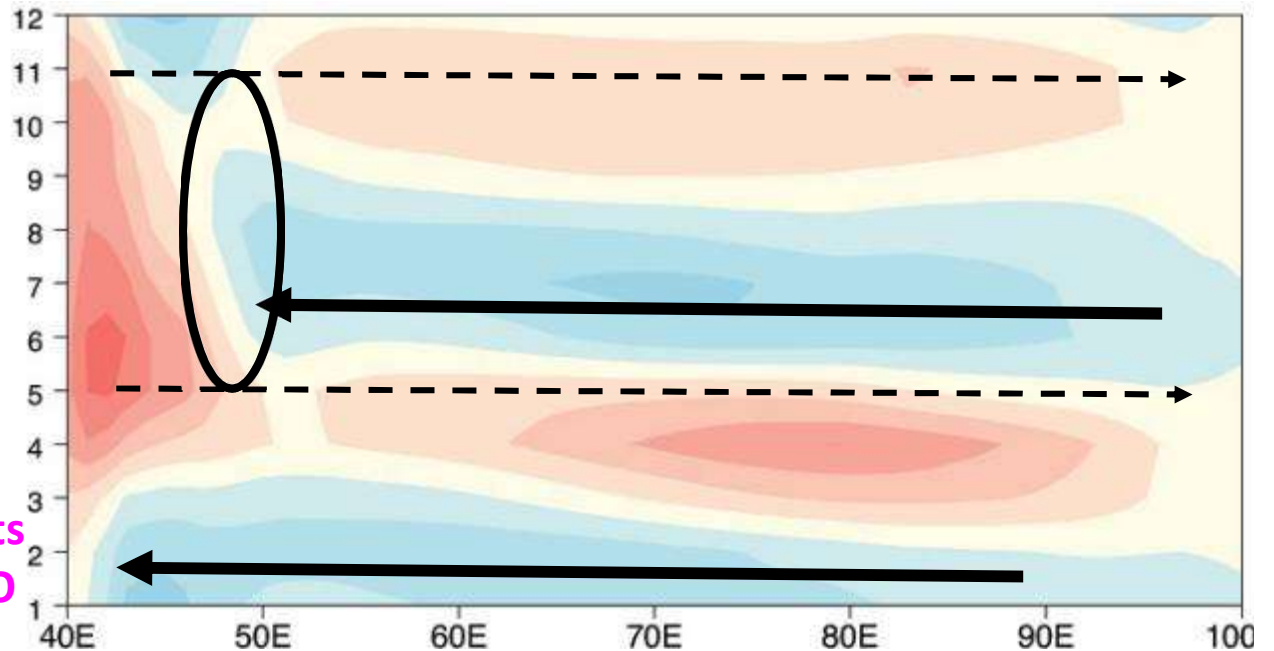
Observations (OSCAR)

Annamalai et al. 2017, JC



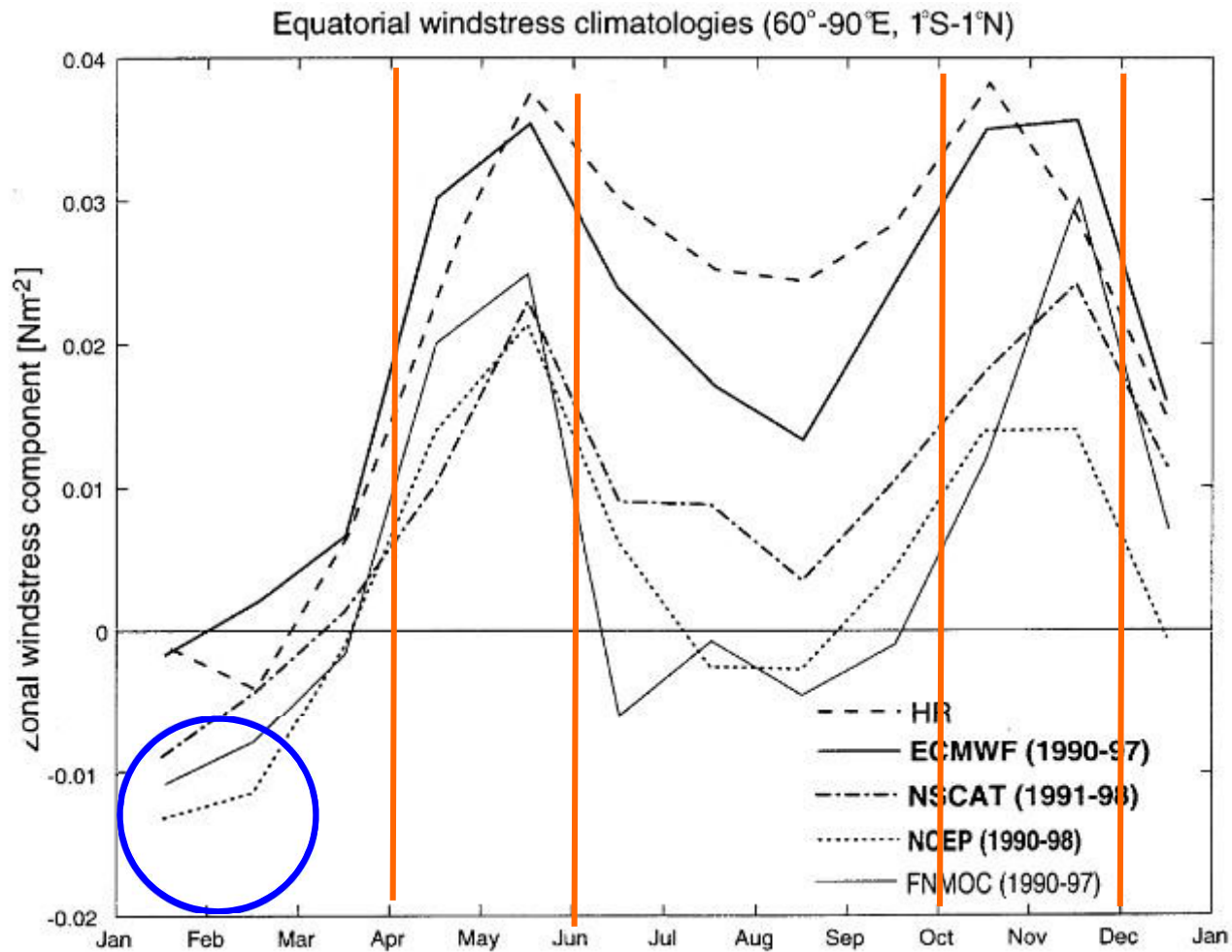
3°S-3°N – surface currents

CMIP5 MMM



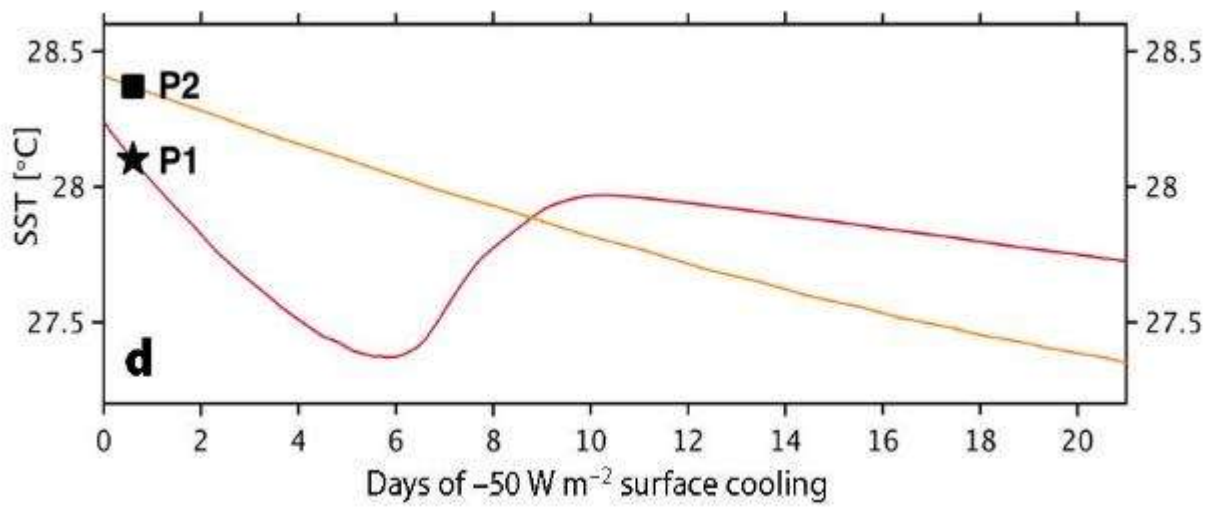
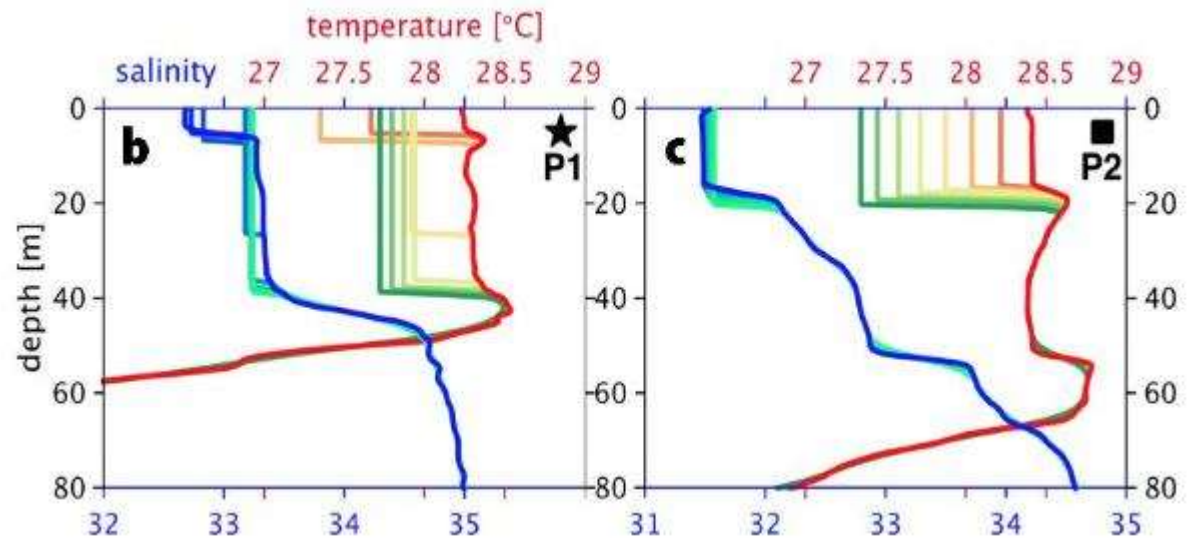
1. Weak eastward WJs
2. Unrealistic westward currents
3. Pile-up of warm waters WEIO





Schott and McCreary (2001)

**Do we need in-situ near-surface wind observations to construct reliable wind-stress climatology?**



As for salinity observations, I think it is good to have moored buoys in addition to Argo floats. The former gives a better time resolution and a robust mean field, and the latter gives a wide spatial coverage. Salinity observations by moored buoys are usually coarse in the vertical direction (typically 20 to 50 m intervals), because many salinity sensors are necessary to increase vertical resolution. Argo floats move vertically, and salinity profiles can be measured by one sensor. So, moored buoys and Argo floats compensate for each other in the sense of vertical resolution too. Moored buoys are costly, and Argo floats are cheap. The basic strategy must be many Argo floats plus several moored buoys at selected sites.

As for Arabian Sea mini warm pool, four moored buoys are necessary at 60, 65, 70, 75E along 5N, which is the region for the thick barrier layer. The spatial structure of the barrier layer is wide in the zonal direction, and thus intervals of 5 degrees should be good enough. South of the southern tip of India (Eq-10N, 70-80E) is the location of sparse Argo float observations, and this region should be filled by deploying more Argo floats. It is perfect if we have data at 10 m intervals from moored buoys and 1 m intervals from Argo floats in the upper 200 m. (I have no idea if observations at each 10 m depth is feasible for moored buoys. But 1-m resolution for Argo floats sounds more likely.)